

Identification of Phase State of Hydrometeors Using Dual-Wavelength Radar

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Abstract

An important objective for the Dual-wavelength Ku-/Ka-band Precipitation Radar (DPR) that will be on board the Global Precipitation Measuring (GPM) core satellite, is to identify the phase state of hydrometeors along the range direction. To assess this, radar signatures are simulated in snow and rain to explore the relation between the differential frequency ratio (DFR), defined as the difference of radar reflectivity factors between Ku- and Ka-bands, and the radar reflectivity factor at Ku-band, Z_{Ku} , for different hydrometeor types. Model simulations indicate that there is clear separation between snow and rain in the Z_{Ku} -DFR plane assuming that the snow follows the Gunn-Marshall size distribution (1958) and rain follows the Marshall-Palmer size distribution (1948). In an effort to verify the simulated results, the data collected by the Airborne Second Generation Precipitation Radar (APR-2) in the Wakasa Bay AMSR-E campaign are employed. Using the signatures of Linear Depolarization Ratio (LDR) at Ku-band, the APR-2 data can be easily divided into the regions of snow, mixed phase and rain for stratiform storms. These results are then superimposed onto the theoretical curves computed from the model in the Z_{Ku} -DFR plane. It has been found that in 90% of the cases, snow and rain can be distinguished if the Ku-band radar reflectivity exceeds 18 dBZ (the minimum detectable level of GPM DPR at Ku-band). This is also the case for snow and mixed-phase hydrometeors. Although snow can be easily distinguished from rain and melting hydrometeors by using Ku- and Ka-band radar, the rain and mixed-phase particles are not always separable. It is concluded that Ku- and Ka-band dual-wavelength radar might provide a potential means to identify the phase state of hydrometeors.

Key word: Air/Space-borne radar, hydrometeor phase, GPM

1. Introduction

The Global Precipitation Measuring (GPM) mission has been proposed for mapping of precipitation globally following the success of the Tropical Rainfall Measuring Mission (TRMM) that measures the precipitation in tropical and sub-tropical region (Simpson et al. 1996; Kummerow et al. 2000; Kozu et al. 2001). The Dual-wavelength Precipitation Radar (DPR) onboard the GPM core satellite is expected to improve accuracy of estimates of precipitation rate and also to enable the retrieval of microphysical properties of hydrometeors, such as particle size distribution and phase state. As the GPM will fly in a higher orbital inclination (65°), the radar observations will be extended to mid- and high-latitude regions where both snow and rain frequently occur. Moreover, because the radar algorithms for estimates of precipitation rate and water content differ for snow and rain, it is necessary to study the feasibility of the GPM DPR for identification of hydrometeor phase states. This capability is also useful in convective rain where a clearly defined bright band is usually absent. Having knowledge of where regions of snow, rain and mixed phase precipitation occur along the radar range direction is important in determining how to allocate estimates of total path attenuation as derived

either by the radiometer or by the use of the radar surface reference technique (Meneghini et al. 2000). Although hydrometeor identification has been studied for polarimetric radar (Ryzhkov and Zrnich 1998; Liu and Chandrasekar 2000; Dolan and Rutledge 2009), such research has not yet been well established using non-polarimetric (co-polarization) radar. Use of radar reflectivity at one frequency alone is not sufficient to distinguish snow from rain echoes because light to moderate rain exhibits a similar range of reflectivities as those of snow. However, the differential frequency ratio (DFR), which is defined as the difference of radar reflectivity factors between Ku and Ka bands, provides useful information that, together with Z_{Ku} , can be useful in improving the capability to distinguish snow and rain.

To explore the capability of the GPM DPR for separation of snow and rain, we start with simulations of radar signatures at the DPR frequencies (Ku- and Ka-band) in snow and rain in an effort to find distinctive characteristics based on assumed particle size distributions. To check the model simulations we use radar measurements taken by the Airborne Second Generation Precipitation Radar (APR-2) during the Wasaka Bay AMSR-E field campaign in 2003 over the sea of Japan on board a NASA P-3 aircraft (Im 2003; Sadowy et al. 2003). The APR-2, operating at

approximately the same frequencies as the DPR, is also capable of obtaining the Linear Depolarization Ratio (LDR) at Ku-band in addition to the measurements of radar reflectivity factors and Doppler velocities at Ku- and Ka-bands. The LDR provides information that can be used to accurately identify the mixed-phase region in stratiform storms, with which the radar data can be easily divided into frozen, melting and liquid regions. The radar measurements from these regions are then compared with the model computations. Thus, the potential to separate hydrometeors among three different phases can be effectively studied.

2. Dual-Wavelength Radar Simulations

Simulation of dual-wavelength radar signatures in snow and rain provides an important means to look into the distinctive features of radar signals from which the phase states of hydrometeors can be possibly identified. What follows is a brief description of the radar parameters involved in this study.

Figures 1 and 2 show the simulated radar signatures in the Z_{Ku} -DFR plane for different values of fixed and variable snow density, respectively. In these simulations the snow size distribution is assumed to be the Gunn-Marshall relation (1958) while the rain is assumed to follow the Marshall-Palmer raindrop size distribution (1948). As shown in the results for the case of fixed snow density (Fig. 1), there is a clear separation between regions of snow (red) and rain (blue) in the Z_{Ku} -DFR space. Departures of the radar signatures from snow to rain grow quickly as the snow density decreases, evidenced by the results of the snow densities with values varying from 0.05 g/cm^3 to 0.5 g/cm^3 ; these separations become increasingly pronounced as the Ku-band radar reflectivity increases. For reference, the curves of constant rain rate (black) are also plotted. For snow these should be interpreted as the equivalent rain rate. In a similar manner, the results of the variable snow density are depicted in Fig.2 where several empirical density-size relations from the results of Wang et al. (2005), Brown and Francis (1995), Heymsfield et al. (2004), Mitchell et al. (1990), and Brandes et al. (2007) are shown. As in the constant snow density case, the results clearly show that the regions of snow and rain can be separated. However, in the variable density case, the separation is much larger than in the case of fixed snow density. The results from Figs.1 and 2 are promising in that they suggest a clear separation between snow and rain returns in the Z_{Ku} -DFR plane. However, these results need to be confirmed by using actual radar measurements. This will be a focus of the remainder of the paper.

3. Radar Phase Identification

To test if the radar measurements can be used to separate the regions of snow, mixed-phase and rain, as indicated by the radar simulations, we first superimpose the radar data onto the theoretical model computations in

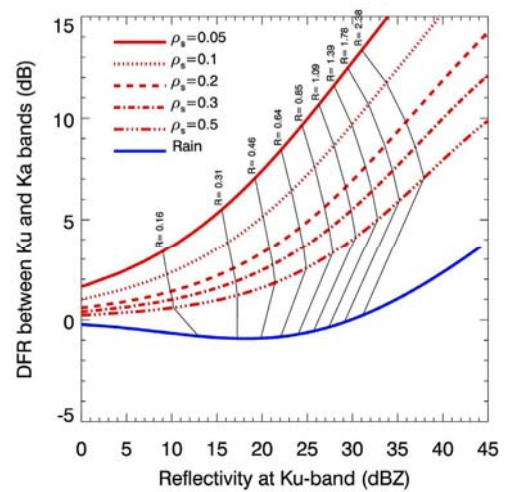


Fig.1 Theoretical relationships between DFR and radar reflectivity factor at Ku band for snow (red curves) and rain (blue curves) as computed from assumed particle size distributions for constant snow density (ρ_s). The black contour lines correspond to equivalent rain rate (R) in mm/h.

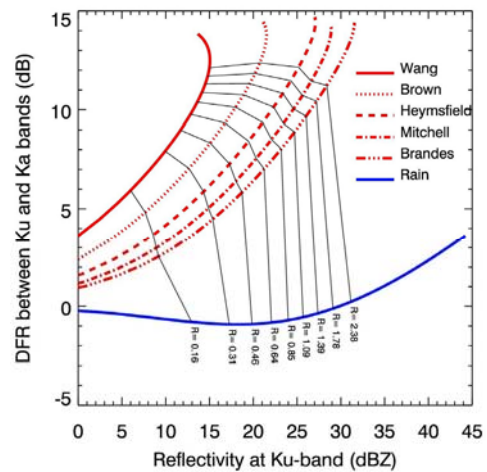


Fig.2 Same as Fig.1 but for snow density that is expressed as a function of particle size. Several empirical density-size relations are used.

the plane of DFR and Ku-band radar reflectivity factor, and then compute two-dimensional probability density function (PDF) from data for quantitative analysis of the results. The data for this study are taken from the APR-2 measurements in stratiform rain on January 23, 2003. Figure 3 exhibits the PDF results in the Z_{Ku} -DFR space for snow (upper cluster) and rain (lower cluster). For the sake of comparison, the contours of the populations are plotted at levels of 40%, 70% and 90%. The theoretical results are also shown in the background with rain given by a heavy-solid line and snow depicted by several thin curves that are associated with different snow densities. Evidently, the majority of the snow and rain echoes are consistent with the model simulations. Note that the results in Fig.3 are derived from all of the measurements of the APR-2 (from 23 equally spaced angle bins), covering a cross-track scan $\pm 25^\circ$ from nadir. In addition,

a simple attenuation correction procedure for rain has been taken in the dataset using the surface reference technique (SRT) (Meneghini et al. 2000). Snow attenuation, on the other hand, can be generally ignored for a downward-looking radar. Attenuation in the melting layer can be corrected based on a melting layer model (e.g., Liao and Meneghini 2005; Liao et al. 2009). However, because the rain height is relatively shallow, with a depth of only about 2 km in this case (not shown), and also because of relatively weak reflectivities attenuation correction, though performed, has little effect on the results. As shown in Fig.3, there are at least 70% cases of these datasets in which the rain and snow are completely separated; for those contours that contain 90% of the data, the rain and snow regions can be fully distinguished if the Ku-band reflectivity exceeds 18 dBZ (which is approximately the DPR minimum detectable level at Ku-band).. Similar comparisons of the PDFs are shown in Fig.4 between the snow and mixed-phase data. It is not difficult to find that 90% of the data from snow and mixed phase are separable. This finding, along with those from Fig.3, suggests that the snow returns usually can be differentiated from rain and mixed phase returns by using measurements of DFR and Ku-band radar reflectivity. In contrast, however, the returns from rain cannot always be distinguished from mixed-phase cases when the Ku-band radar reflectivity factors are in the range from 25 to 35 dBZ, as illustrated in Fig.5. This overlap region is associated with the trailing part of the melting layer in which rain and melting snow returns yield approximately the same magnitudes of DFR and radar reflectivity factor at Ku-band. For values of Z_{Ku} below 25 dB, the returns are exclusively associated with rain while above 35 dB, the returns are primarily associated with the mixed phase region. These findings are supported by the similar comparisons (not shown) from data acquired from other flights on different days. It is worth noting that the comparison of the radar measurements to the model simulations in snow for the fixed and variable snow densities reveals that the radar simulations are in better agreement if the averaged bulk (or fixed) snow density is used.

4. Summary

To develop an algorithm for the Ku- and Ka-band dual-wavelength spaceborne radar for the identification of hydrometeor phase state, such as snow, melting snow (mixed-phase) and water, we start with model simulations that are focused on snow and rain. As the rain and snow particle size distributions can be described by an exponential function, i.e. the Marshall-Palmer size distribution for rain and the Gunn-Marshall size distribution for snow, theoretical computations of the relations between DFR and radar reflectivity factor at Ku-band are performed. It is shown that snow and rain can usually be clearly distinguished in the Z_{Ku} -DFR plane. The data taken from the NASA JPL APR-2 Ku- and Ka-band airborne radar during Wakasa Bay experiment are employed to check the model simulations

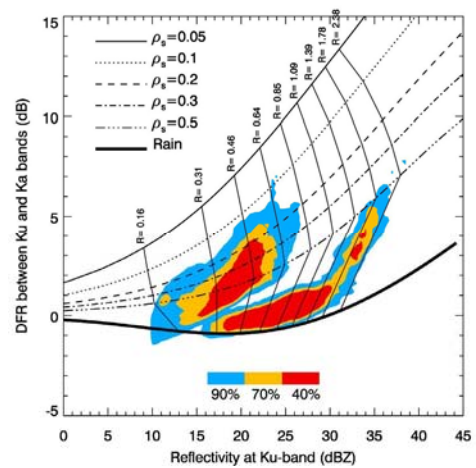


Fig.3 Two-dimensional PDF for snow (upper cluster) and rain (lower cluster) as the contours of data populations are plotted at 40%, 70% and 90%.

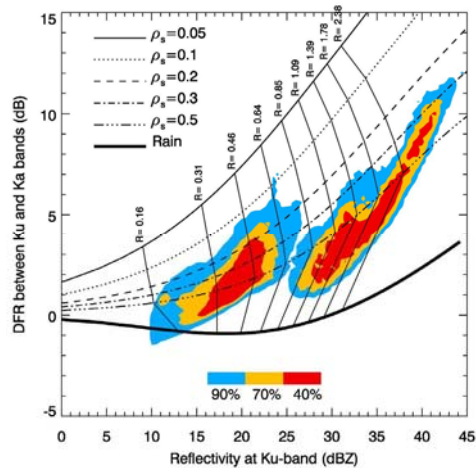


Fig.4 Two-dimensional PDF for snow (left cluster) and mixed-phase (right cluster) as the contours of data populations are plotted at 40%, 70% and 90%.

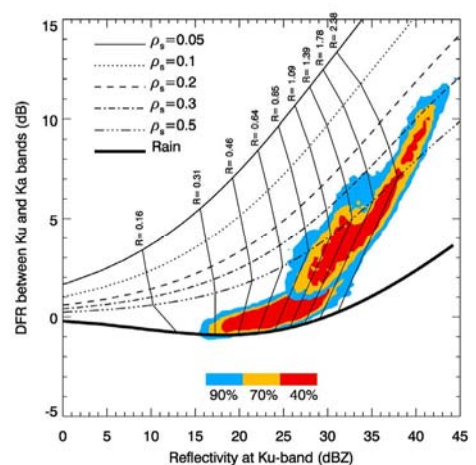


Fig.5 Two-dimensional PDF for rain (lower-left cluster) and mixed-phase (upper-right cluster) as the contours of data populations are plotted at 40%, 70% and 90%.

and also to assess the feasibility of the dual-wavelength radar technique for identification of hydrometeor phase. Having superimposed the data collected in stratiform rain on January 23, 2003 onto the theoretical curves in the plane of DFR and radar reflectivity at Ku-band, it is found that the radar measurements are generally consistent with the model simulations for snow and rain. When compared with measurements, better agreement is obtained using fixed snow density than variable snow density. Analysis of the two-dimensional PDF of the data from snow, rain and mixed-phase indicates that snow can be easily distinguished from rain and mixed-phase hydrometeors if the radar reflectivity factor at Ku-band is greater than 18 dBZ. Rain, however, is not always separable from mixed-phase data. In stratiform rain, the region of overlap occurs at ranges near the lower boundary of the melting layer where the snow melting is nearly complete and where the radar reflectivities and DFR values from the melting snow resemble those from rain. Apart from this overlap region ($25 \text{ dBZ} < Z_{\text{Ku}} < 35 \text{ dBZ}$), the separation between the regions of rain and mixed phases appears feasible.

Although the radar technique described in this study presents a promising way to separate the regions of snow, rain and mixed-phase hydrometeors, attenuation due to rain and mixed-phase particles may complicate the identification because Ka-band experiences much more attenuation than the Ku-band in general, resulting in an increase in DFR. Without proper attenuation correction, it is possible to misclassify the hydrometeors. One way to circumvent this problem is to use the SRT and dual-wavelength radar retrieval algorithms. The SRT provides an effective means for attenuation correction at the surface based on the difference of surface returns between rain and rain-free areas. The dual-wavelength radar backward approach which works on gate-by-gate basis can be possibly incorporated with the phase identification scheme to identify the phase states beginning from the surface and moving upward toward the storm top (Liao and Meneghini 2005; Liao et al. 2008). To develop a fully functional radar algorithm, an investigation into an optimal selection of the DFR-Z(Ku) relation that effectively separates different phase regions is needed. To achieve this, a further study will be undertaken, which includes an analysis of dual-wavelength radar measurements and the model computations linking DFR to radar reflectivity under various microphysical conditions, and an exploration of convective storm in which an obvious bright-band is absent. The development of an effective dual-wavelength method to distinguish the rain, snow and mixed-phase hydrometeors will be an important step toward an accurate, efficient DPR profiling algorithm.

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